

1 **Fluid-driven fracture propagation in heterogeneous media:**
2 **probability distributions of fracture trajectories**

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9 **Abstract**

10 Hydraulic fracture trajectories in rocks and other materials are highly affected by spatial hetero-
11 geneity in their mechanical properties. Understanding the complexity and structure of fluid-driven
12 fractures and their deviation from the predictions of homogenized theories is a practical problem
13 in engineering and geoscience. We conduct a Monte Carlo simulation study to characterize the
14 influence of heterogeneous mechanical properties on the trajectories of hydraulic fractures propa-
15 gating in elastic media. We generate a large number of random fields of mechanical properties and
16 simulate pressure-driven fracture propagation using a phase-field model. We model the mechanical
17 response of the material as that of an elastic isotropic material with heterogeneous Young modulus
18 and Griffith energy release rate, assuming that fractures propagate in the toughness dominated
19 regime. Our study shows that the variance and the spatial covariance of the mechanical properties
20 are controlling factors in the tortuousness of the fracture paths. We characterize the deviation of
21 fracture paths from the homogenous case statistically, and conclude that the maximum deviation
22 grows linearly with the distance from the injection point. Additionally, fracture path deviations
23 seem to be normally distributed, suggesting that fracture propagation in the toughness-dominated
24 regime may be described as a random walk.

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25 I. INTRODUCTION

26 Hydraulic fracturing process is widely present in geomechanics. A natural example is
27 the fracture generated by the migration of magma through the lithosphere, resulting in thin
28 dikes and sills which extend over several kilometers [1, 2]. Some man-made examples arise
29 from the engineering field, as for instance the geologic sequestration of carbon dioxide [3],
30 the stimulation of geothermal reservoirs to increase heat extraction [4, 5], or the enhanced
31 oil and gas extraction from unconventional low-permeability reservoirs [6, 7]. The latter
32 two applications have a growing interest which arise from environmental concerns. Heat
33 extraction from geothermal reservoirs has the potential to become a crucial source of re-
34 newable carbon-free energy [8]. Unconventional oil and gas are expected to boost energy
35 supply by 2035 [9], and technically recoverable unconventional gas resources world-wide from
36 low-permeability reservoirs are roughly estimated around 8,000 trillion cubic feet [10].

37 One of the major concerns of the application of hydraulic fracturing technologies is the
38 groundwater contamination. Created fractures may unintentionally connect the reservoir
39 and the surrounding environment, providing hydraulic access between both and leading to
40 leakage of fracturing fluid or gas [11, 12]. Fracture propagation has a stochastic component
41 which results into complex trajectories. Experiments have shown that fracture trajectories
42 are highly affected by the *in situ* stress [13, 14], heterogeneity of the reservoir formations [15,
43 16], or the presence of natural fractures [17], among other factors. These effects have also
44 been studied through complex numerical simulations [18–20].

45 Estimating the uncertainties in fracture propagation through numerical simulations is a
46 challenging task. Monte Carlo simulation method is the simplest and widest-used approach
47 to estimate uncertainty [21]. The biggest disadvantage is that a large number of realizations
48 are required. The influence of the input parameters on the model output can be quantified
49 in the local or global context. The former keeps fixed all the inputs while varying one of
50 them. The latter conducts sensibility analyses through, for instance, stochastic multiscale
51 methods [22, 23]. Another approach is the stochastic finite element method, but unfortu-
52 nately its applications are mostly limited to linear problems [24], although some attempts
53 to extend the method to nonlinear problem have been performed [25, 26]. Some attempts
54 to quantify the uncertainty associated to fracture process due to heterogeneous material
55 properties have been conducted within the Monte Carlo simulation framework [21, 27].

56 The estimate of uncertainty of fluid-driven fracture paths through the Monte Carlo simu-

57 lation framework requires both a random field generator algorithm of mechanical properties,
58 and a mathematical model of the fracture process. Techniques for generating random fields
59 can be broadly grouped into discrete and continuous models [28]. The former generate
60 random fields of discrete properties, as for instance, the number of fractures in fracture net-
61 works. On the other hand, continuous models describe properties that vary continuously, as
62 for instance, the Young modulus or the hydraulic conductivity. The most common contin-
63 uous models are [29]: the moving average methods, the discrete Fourier transform (DFT)
64 method, the covariance matrix decomposition, the fast Fourier transform (FFT) method,
65 the turning bands method, and the local average subdivision method.

66 Fluid-driven fracture propagation in elastic media can be simulated through analytical
67 and numerical models. Analytical solutions are available for simple injection protocols and
68 fracture geometries. [30] provides a review of the state-of-the-art for the penny-shape fracture
69 propagation. More complex geometries or injection protocols require numerical models.

70 Numerical models for fluid-driven fracturing can be classified into discrete or continuous
71 approaches. In the discrete approaches, fractures are simulated as discontinuities. Some
72 discrete models are implemented using the finite element method (FEM), where the prop-
73 agation of the fractures may be simulated by splitting nodes [31], breaking elements [32],
74 enriching the displacement field with new degrees of freedom [33, 34] or using cohesive inter-
75 face elements [35]. Other models are the spring network approach, where fractures grow by
76 splitting nodes and reconnecting springs [36], or the mesh free methods including the crack-
77 ing particle method [37, 38], the immersed particle method [39], or the discrete element
78 approach [40, 41].

79 Continuous approaches describe both the intact solid and fracture domains as a unique
80 continuum. The geometric simplicity and the topology of the discretization is held during
81 the fracture propagation, which makes these models computationally appealing. Examples
82 of continuous approaches include peridynamics, in which the solid is assumed to be com-
83 posed of material points that interact with all its neighbors within a nonlocal region called
84 horizon [42–45], or the screen-Poisson method where fractures are regularized through a
85 screened Poisson equation [46].

86 A novel continuous mathematical approach for fluid-driven fracture propagation in brittle
87 elastic media is the phase-field model. It is based on the Griffith interpretation to model
88 quasi-static brittle fracture propagation [47]. Fractures are treated as a diffuse interface
89 defined by the phase-field variable that interpolates between the broken and unbroken re-

90 gions. This approach enjoys the following advantages [48]: (1) the approach is fixed-mesh,
91 (2) the formulation is based on energy minimization and, consequently, nucleation, propa-
92 gation, joining, or branching are automatically computed without any additional handling,
93 and (3) fracture propagation in heterogeneous media is simulated without any additional
94 modification in the computational framework. Moreover, its capability of simulating fluid-
95 driven fracturing in an elastic medium has been assessed against analytical solutions with
96 encouraging results [49], and compared with laboratory experiments with promising perfor-
97 mance [50].

98 Here, we study the uncertainty of fluid-driven fracture trajectories associated to hetero-
99 geneity of rock mechanical properties. We adopt a Monte Carlo simulation framework where
100 random fields of several mechanical properties are generated through a FFT method. We
101 quantify the influence of the mechanical parameters on the fracture path in the local context.
102 Moreover, fracture trajectories are simulated using a phase-field approach. For simplicity,
103 we restrict our study to a 2-dimensional (2-D) elastic case study under plain strain condi-
104 tions, infinitesimal deformations, and isothermal conditions. We assume fractures propagate
105 under the toughness dominated regime with no-fluid lag. The main physical consequence of
106 these assumptions is the restriction of our conclusions to very low permeable rocks where a
107 low-viscous fluid is injected. This situation is quite real since hydraulic fracturing is a tech-
108 nique for enhancing oil and gas production in formations with low permeability. Moreover,
109 fracturing jobs are more efficient when fractures propagate under the toughness dominated
110 regime since the energy expended in fracturing the rock is much larger than the viscous
111 dissipation [51]. The no-fluid lag assumption is valid when the confining stress is suffi-
112 cient large [52], although the literature about modeling hydraulic fracturing with fluid-lag is
113 scarce [53, 54]. The simulated trajectories allow us to characterize and statistically quantify
114 the influence of several statistical parameters of the random fields on the tortuousness in
115 fracture trajectories. The research question is, then, how rock heterogeneity makes the frac-
116 ture trajectory deviate from the deterministic path, whether the deviations follows random
117 walks or not, and how far fractures deviate.

118 The paper is organized as follows. Section 2 provides the governing equations of the
119 Monte Carlo simulation framework. First, we introduce the random field generator algo-
120 rithm of mechanical properties of the elastic medium. Afterward, the fluid-driven fracture
121 model based on the phase-field approach is presented. A case study where we apply the
122 methodology is described in section 3. The results of the numerical simulations and their

123 analyses are reported in section 4. Finally, some conclusions are drawn in section 5.

124 II. METHODOLOGY

125 In this section, we begin by describing the geometric representation of the fracture/matrix
126 domain. We then introduce the methodology used to generate correlated synthetic random
127 fields of Young modulus and Griffith energy release rate through a Fourier space method.
128 Finally, we describe the phase-field approach for the simulation of quasi-static, brittle, fluid-
129 driven fracture propagation in elastic solids.

130 II.1. Geometry

131 Consider a domain $\Omega \subset \mathbb{R}^\delta$ with spatial dimension $\delta \in \{2, 3\}$. The problem domain
132 comprises an impermeable elastic subdomain, Ω_E , and a pressurized fracture subdomain, Ω_F ;
133 *i.e.*, $\Omega = \Omega_E \cup \Omega_F$ (Fig. 2). A slightly compressible Newtonian fluid is injected into the
134 fracture at known pressure. While no mass exchange occurs between subdomains, the fluid
135 pressure exerts a force over the elastic subdomain that induces the propagation on the
136 fracture. The boundary of Ω_i is denoted by $\partial\Omega_i$, where $i = E, F$. The boundary of Ω_E
137 is divided into two subsets, $\partial_D\Omega_E$ and $\partial_N\Omega_E$, where Dirichlet and Neumann boundary
138 conditions are imposed, respectively.

139 II.2. Random field generator

140 We assume that the heterogeneous mechanical properties of real materials can be de-
141 scribed, at the macroscopic scale, as a random yet continuous spatial variation of these
142 quantities. The random field theory [55], through the design of suitable random field gen-
143 erators, is a common tool for modeling this spatial variability. The cornerstone of this
144 approach is the selection of a probability distribution that matches the natural distribution
145 of the property concerned.

146 Numerous field data analyses and experimental studies have been conducted to char-
147 acterize the probability distribution of constitutive properties in natural and engineered
148 materials. The hydraulic transmissivity, conductivity, and storage coefficient of aquifers
149 have been found to follow lognormal distributions [56]. Mechanical properties are often

150 found to be lognormally distributed, from the elastic modulus and strength of concrete [57],
 151 to the geomechanical parameters of rocks [58] and the elastic modulus of the cytoskeletal
 152 network of living cells [59].

153 Given the ubiquitous presence of lognormal properties in nature, the lognormal distri-
 154 bution has been adopted in many studies to describe material properties in synthetic mod-
 155 els. Some examples include the hydraulic properties of soil in probabilistic slope stability
 156 analysis [60], infiltration and unsaturated flow [61], flow through porous rocks [62, 63] or
 157 macrodispersion in aquifers [64]. It has also been adopted to describe the mechanical prop-
 158 erties of steel in fatigue and fracture analyses of bridges [65], or the cohesion and friction
 159 angle of rocks in the design of rock slopes [66, 67].

160 Here, we consider both the Young modulus, E , and the Griffith energy release rate, G_c ,
 161 to be isotropic and stationary lognormal fields. We do not consider spatial variations in the
 162 Poisson ratio, ν , which is rather taken as a deterministic constant throughout the domain.
 163 The assumption of stationarity is appropriate when a single material is modeled, and it
 164 implies that all values of a given mechanical property at different points of the solid are
 165 drawn from the same distribution with the same statistical parameters.

166 Given a random field H in a spatial domain D , $\{H(\mathbf{x}); \mathbf{x} \in \mathbf{D} \subseteq \mathfrak{R}^d\}$, following a lognor-
 167 mal distribution with mean μ_H and standard deviation σ_H , its logarithm, $\ln H$, is normally
 168 distributed with mean, $\mu_{\ln H}$, and standard deviation, $\sigma_{\ln H}$, given by:

$$169 \quad \sigma_{\ln H}^2 = \ln \left(1 + \sigma_H^2 / \mu_H^2 \right), \quad (1)$$

170 and

$$171 \quad \mu_{\ln H} = \ln(\mu_H) - \frac{1}{2} \sigma_{\ln H}^2. \quad (2)$$

172 We assume that values of H at two points in the domain show correlation that decreases with
 173 the distance between them. Correlation is modeled through the so-called spatial correlation
 174 function. Here we adopt an isotropic Gaussian type function, $\rho(r)$, which is commonly used
 175 to describe geologic materials [68]:

$$176 \quad \rho(r) = \sigma_{\ln H}^2 \exp \left(-r / (\sqrt{2} l_c) \right)^2, \quad (3)$$

177 where r is the spatial distance and l_c is the correlation length. Points separated by distances
 178 larger than l_c are weakly correlated and therefore essentially independent. We assume that
 179 both fields, E and G_c , have the same correlation function. The field $\ln H$ is expressed as:

$$180 \quad \ln H(\mathbf{x}) = \mu_{\ln H} + \sigma_{\ln H} Z(\mathbf{x}), \quad (4)$$

181 where $Z(\mathbf{x})$ follows a standard normal distribution with spatial autocorrelation $\rho(r) =$
 182 $\exp(-r/(\sqrt{2}l_c))$. We generate random Gaussian fields efficiently through a spectral ap-
 183 proach based on the Fast Fourier Transform (FFT) method [69]. Random fields in the
 184 spatial domain, $Z(\mathbf{x})$, are obtained as the inverse Fourier transform of independent random
 185 Fourier increments which are generated on a wave number grid [70].

186 Regarding the dependence between the $E(\mathbf{x})$ and $G_c(\mathbf{x})$ fields, we assume that at a given
 187 location \mathbf{x}_i both lognormal fields depend on each other through only the local variable $Z(\mathbf{x}_i)$
 188 as follows:

$$189 \quad \ln E(\mathbf{x}_i) = \mu_{\ln E} + \sigma_{\ln E} Z(\mathbf{x}_i), \quad (5)$$

$$190 \quad \ln G_c(\mathbf{x}_i) = \mu_{\ln G_c} + \sigma_{\ln G_c} Z(\mathbf{x}_i), \quad (6)$$

192 or, alternatively,

$$193 \quad \ln E(\mathbf{x}_i) = a \ln G_c(\mathbf{x}_i) + b, \quad (7)$$

194 where the constants a and b are:

$$195 \quad a = \frac{\sigma_{\ln E}}{\sigma_{\ln G_c}}, \quad (8)$$

$$196 \quad b = \frac{\mu_{\ln E} \sigma_{\ln G_c} - \mu_{\ln G_c} \sigma_{\ln E}}{\sigma_{\ln G_c}}. \quad (9)$$

198 The above assumption implies that the correlation coefficient between both variables is one.
 199 It is motivated by the fact that in the same material the elastic modulus and the strength
 200 are closely related.

201 **II.3. Rock damage model**

202 We model the propagation of fluid-driven fractures through a quasi-static phase-field
 203 formulation. A detailed description of the model formulation can be found in [49]. Fractures
 204 are simulated as diffuse interfaces defined by the phase-field variable, $d \in [0, 1]$, which
 205 can be interpreted as a damage indicator. We define $d \rightarrow 1$ at the fracture and $d=0$ at
 206 the undamaged regions. In the following formulation, we adopt an anisotropic damage
 207 formulation [71] which allows for fracturing in tension only. The anisotropic formulation
 208 provides realistic fracture evolution patterns for rocks and other materials with asymmetric
 209 response for fracturing under tension and compression. The equilibrium equations for the
 210 damage model are given by:

211
$$-\nabla \cdot \{[g(d)+\kappa] \sigma_0^+ + \sigma_0^-\} - p_f \nabla g(d) - \bar{\mathbf{f}} = \mathbf{0}, \quad \text{in } \Omega, \quad (10)$$

212

213
$$\sigma \bar{\mathbf{n}} = \bar{\mathbf{t}}, \quad \text{in } \partial_N \Omega. \quad (11)$$

214 where σ_0^\pm refers to Cauchy tensile/compressive stresses in the undamaged solid, $g(d)$ is a
 215 degradation function, κ is a small positive parameter, p_f is the pressure of the fluid inside
 216 the fracture, $\bar{\mathbf{f}}$ is the body force vector per unit volume, $\bar{\mathbf{t}}$ is the vector of applied traction
 217 surface forces, and $\bar{\mathbf{n}}$ is the outward normal unit vector. The small positive parameter $\kappa \approx 0$
 218 prevents numerical singularities and is chosen as small as possible such that the algebraic
 219 conditioning number of the numerical method guarantees well-posedness for partly-broken
 220 systems [72].

221 The tensile/compressive stress σ_0^\pm is given by:

222
$$\sigma_0^\pm(\mathbf{u}) = \sum_{a=1}^{\delta} \left[\lambda \left\langle \sum_{i=1}^{\delta} \varepsilon_i \right\rangle_{\pm} + 2\mu \langle \varepsilon_a \rangle_{\pm} \right] \mathbf{n}_a \otimes \mathbf{n}_a, \quad (12)$$

223 where $\langle x \rangle_{\pm} = (x \pm |x|)/2$, λ and μ are the Lamé constants, δ is the number of space
 224 dimensions, ε_a is the a^{th} principal strain, and \mathbf{n}_a is the principal strain direction associated
 225 to the principal strain ε_a . The eigenvalue bases $\mathbf{n}_a \otimes \mathbf{n}_a$ are computed in terms of the
 226 principal strains ε_i and the strain tensor ε as follows [73] :

227
$$\mathbf{n}_a \otimes \mathbf{n}_a = \frac{1}{\prod_{b \neq a}^{\delta} (\varepsilon_a - \varepsilon_b)} \prod_{b \neq a}^{\delta} (\varepsilon - \varepsilon_b \mathbf{1}). \quad (13)$$

228 We employ a quadratic degradation function $g(d) = (1-d)^2$ [74]. While other degradation
 229 functions have been proposed in the literature, the impact of its particular functional form
 230 diminishes after the fracture has formed [75], and Γ -convergence has so far only been proved
 231 for the quadratic polynomial [76].

232 The strong form of the phase-field problem reads:

233
$$\frac{g_c}{\ell} (d - \ell^2 \nabla^2 d) = 2(1-d)H^+, \quad \text{in } \Omega, \quad (14)$$

234

235
$$\nabla d \cdot \bar{\mathbf{n}} = 0, \quad \text{in } \partial_N \Omega, \quad (15)$$

236 where ℓ is the length scale parameter, and H^+ is the maximum local history field. The
 237 characteristic length ℓ controls the band width across which the regularization of the frac-
 238 ture surface takes place. Phenomena with characteristic length scales smaller than ℓ cannot

239 be captured by our model, e.g. the observed self affine scaling of fracture roughness [77].
 240 Numerically, ℓ has a slight influence over the critical load that triggers unstable fracture
 241 propagation—the load decreases as ℓ increases [50, 74]. The memory term H^+ is the maxi-
 242 mum historic energy generated by tensile strain, and is given by:

$$243 \quad H^+(\mathbf{u}, p_f, t) = \max_{s \in [0, t]} (\psi_0^{e+}(\varepsilon) + p_f \nabla \cdot \mathbf{u} + \mathbf{u} \cdot \nabla p_f), \quad (16)$$

244 where \mathbf{u} is the displacement field, t is time, and ψ_0^{e+} is the positive undamaged elastic
 245 energy density. More generally, $\psi_0^{e\pm}$ splits the total energy among tension and compression
 246 as follows:

$$247 \quad \psi_0^{e\pm}(\varepsilon) = \frac{\lambda}{2} \left(\left\langle \sum_{i=1}^{\delta} \varepsilon_i \right\rangle_{\pm} \right)^2 + \mu \sum_{i=1}^{\delta} (\langle \varepsilon_i \rangle_{\pm})^2, \quad (17)$$

248 and fulfills $\psi_0^e(\varepsilon) = \psi_0^{e+}(\varepsilon) + \psi_0^{e-}(\varepsilon)$. Irreversibility is enforced in the model through H^+ ,
 249 which ensures that damaged regions do not heal, even when stress disappears [71].

250 We consider that fractures propagate under the toughness-dominated regime, *i.e.* the
 251 energy expended in fracturing the rock is much larger than the viscous dissipation due to
 252 the movement of the fluid inside the fracture [78]. We further assume that the fluid viscosity
 253 is negligible, and as a result fluid pressure inside the fracture is uniform and equal to the
 254 prescribed boundary pressure. Consequently, since the pressure inside the fracture is uniform
 255 and its value is known at the injection point, there is no need to solve the flow equation
 256 along the fracture to update p_f . Since the fracture domain is modeled as a field, the fracture
 257 path has to be inferred from the phase-field. In our 2D setting, the trajectory is the line
 258 where the gradient of the phase-field is zero and the phase-field is above a threshold d_f (here
 259 we adopt $d_f = 0.9$ [49])-

260 The coupled problem—*i.e.* the equilibrium equation (10) together with the damage
 261 model (14)—is implemented using the finite element software COMSOL Multiphysics [79].
 262 The integration in time is fully implicit and adaptive, and it is based on backward differen-
 263 tiation formulae [80].

264 III. CASE STUDY

265 We simulate the propagation of a single fracture under the toughness dominated regime
 266 in an elastic solid whose Young modulus and Griffith critical energy release rate spatially
 267 vary following stationary isotropic log-normal fields (Fig. 1). Since our main purpose is to

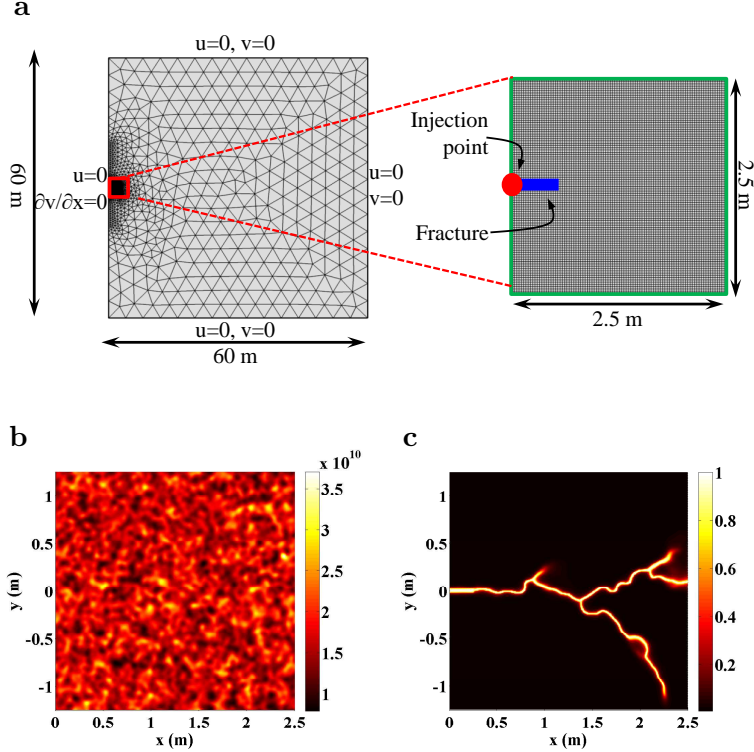


FIG. 1. Schematic description of the case study: problem geometry, computational mesh and illustration of the relevant fields. (a) We simulate fluid injection at prescribed pressure, and study the propagation of an initially-horizontal fracture of length 0.2 m. The outer domain boundaries are fixed except at the left side, where we impose symmetric displacements. The computational mesh is structured in the inner domain, with 4-noded quadrilateral elements, and unstructured in the outer domain with 3-noded triangular elements. Maximum mesh sizes in the inner and outer domains are, respectively, 8 mm and 4 m. (b) A realization of lognormal Young modulus with Gaussian spatial correlation function. The mean is μ_E is 17 GPa, σ_E is 3.4 GPa, μ_{g_c} is 100 Pa·m, σ_{g_c} is 20 Pa·m, and the correlation length is l_c is 0.10 m. (c) Using the previous realization, we simulate the fracture propagation. Here we show a map of the phase-field variable at the end of the simulation.

268 investigate the effect of the spatial variability of E and G_c , we consider ν and the boundary
 269 conditions to be deterministic.

270 The domain is a square of side 60 m, and an initial horizontal fracture of length 0.2 m is
 271 located at the center of the left side, Fig. 1(a). We inject fluid at prescribed pressure through
 272 a point located at (0, 0) in the left side of the fracture. The boundary conditions are zero
 273 displacements, except at the left side where symmetry boundary conditions are imposed.
 274 We assume plane strain conditions.

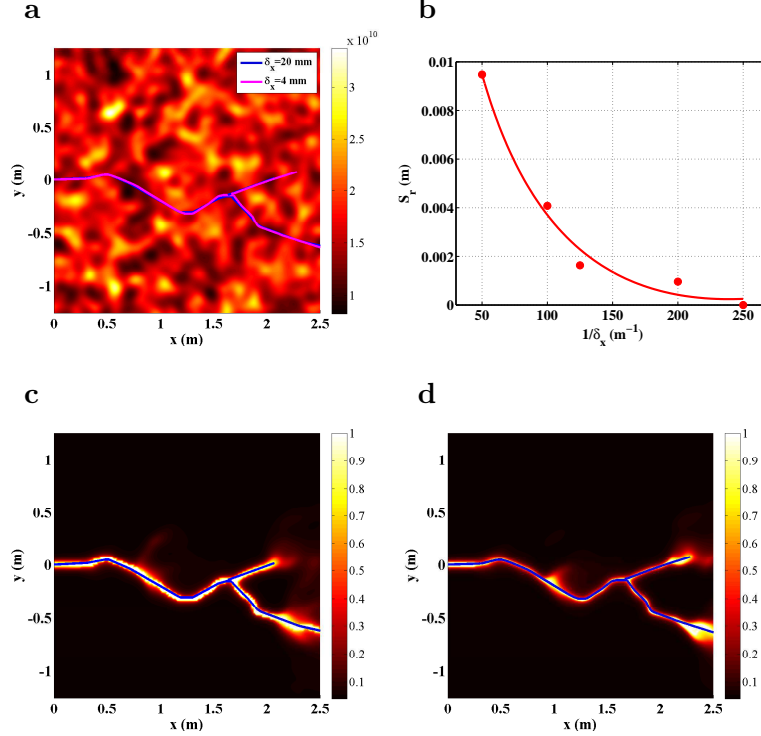


FIG. 2. Mesh convergence study of the fracture trajectories. (a) We plot one realization of 2D lognormal Young modulus with a Gaussian correlation function. μ_E is 17 GPa, σ_E is 3.4 GPa, μ_{g_c} is 100 Pa·m, σ_{g_c} is 20 Pa·m, and l_c is 0.10 m. The fracture paths computed with inner mesh sizes of 20 and 4 mm are depicted in blue and magenta respectively. (b) Here we depict the value of S_r of the fracture path for every simulation, taking as reference solution the simulation with the smallest size. The solution converges to the reference case as the mesh size is smaller. We depict the phase-field variable calculated with an inner mesh size of 20 and 4 mm and the fracture paths in blue in the panels (c) and (d) respectively. As the mesh size is smaller, the 2-D damaged area shrinks and converges to the fracture path.

275 We minimize boundary effects by studying the fracture propagation inside an inner square
 276 of side 2.5 m—red square of Fig. 1(a)—where we use a highly refined mesh. The outer
 277 domain is discretized as an unstructured mesh with linear triangular elements of maximum
 278 size 4 m, while the inner domain comprises 4-noded quadrilateral elements whose size is
 279 selected through a convergence study.

280 Given one realization of E and G_c , we simulate the propagation of the initial "seed"
 281 fracture with mesh sizes ranging from 2×10^{-2} m to 4×10^{-3} m, Fig 2(a). We then compute
 282 the root-mean-square deviation, S_r , of the simulated fracture paths with respect to the
 283 result provided by the simulation with the smallest mesh size, Fig 2(b). The damaged area

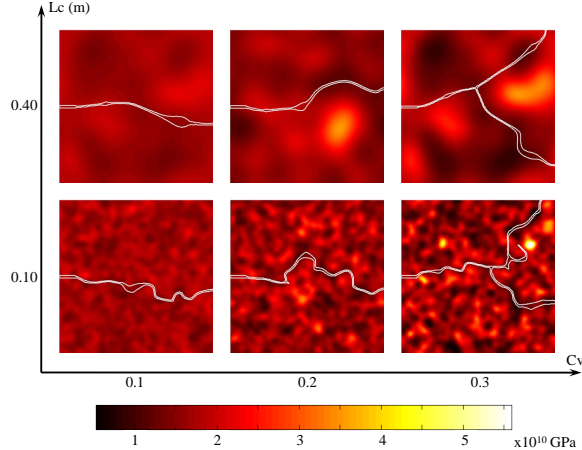


FIG. 3. Schematic description of the influence of C_v and l_c . We plot one realization of the Young modulus for some values of C_v and l_c . In the same panels we also plot in white the fracture paths. It can be appreciated how the trajectories become more tortuous as either the variability of the Young modulus or the correlation length increase. The fracture propagates following the weakest regions, which in our simulations also have the lowest values of the Young modulus.

284 provided by the phase-field variable has the same dimension as the domain, but it shrinks
 285 and converges to the fracture path as the mesh size is smaller, Fig 2(c)-(d). The values of S_r
 286 for the considered sizes suggest that a good choice is $\delta_x = \delta_y = 8$ mm, which provides results
 287 that are very similar to the simulations with the smallest size. As a reference, the mesh
 288 shown in Fig 2(a) is composed of 15,161 triangular elements and 97,969 square elements.

289 In the simulations analyzed in the next section, the mechanical properties of the elas-
 290 tic solid are: mean Young modulus $\mu_E = 17$ GPa, mean Griffith critical energy release
 291 rate $\mu_{G_c} = 100$ Pa·m, Poisson ratio $\nu = 0.2$, and length scale parameter $\ell = 4 \times 10^{-2}$ m.
 292 We depict one realization of the Young modulus field in Fig. 1(b), where σ_E is 3.4 GPa, σ_{g_c}
 293 is 20 Pa·m, and the correlation length is 0.10 m. The injection pressure varies from 1 MPa
 294 at time zero to 4 MPa at time 7.2×10^3 s. We stop the simulation when the fracture reaches
 295 any of the sides of the inner domain. The key result of the simulation is the phase-field
 296 variable, which provides the fracture path. As an example of a typical fracture evolution,
 297 we show the phase field variable at the end of a simulation in Fig. 1(c), which corresponds
 298 to the realization of Young modulus field shown in Fig. 1(b).

299 **IV. NUMERICAL RESULTS AND DISCUSSION**

300 We study the influence of the statistical properties of the heterogeneous fields E and G_c
 301 on the fracture trajectories. These trajectories can be understood as perturbations of the
 302 homogeneous case—with uniform mean values of E and G_c —for which fractures propagate
 303 along a straight path. Spatially variable mechanical properties make fractures deviate from
 304 the straight trajectory and even branch, corresponding to random propagation.

305 We illustrate several random fracture trajectories in Fig. 3, where we plot one realization
 306 of the Young modulus for some values of the coefficient of variation, $C_v = \sigma/\mu$, and the
 307 correlation length, l_c . In the same panels we also plot the simulated fracture paths in white.
 308 It can be appreciated how the variability of the fields increases with C_v , how the size of the
 309 regions with a similar magnitude of the field becomes larger with l_c , and how the trajectories
 310 become more tortuous as either C_v increases or l_c decreases. The strongest regions are those
 311 with the highest value of E , and are coincident with the largest G_c . The fractures propagate
 312 following the weakest regions, depicted in dark colors, and try to avoid the strongest ones.
 313 This result is in agreement with some experiments conducted in concrete with different
 314 aggregate strengths [81].

315 We analyze the effects of C_v and l_c on the fracture path in the local context, *i.e.* we fix
 316 one of the parameters while varying the other one. We consider seven data sets of statistical
 317 parameters, listed in Table I, where sets correspond to three values of $C_v = \{0.1, 0.2, 0.3\}$ for
 318 constant $l_c = 0.10$ m, and five values of $l_c = \{0.05, 0.10, 0.15, 0.20, 0.40\}$ m for constant $C_v =$
 319 0.2. The mean value of the Young modulus, μ_E , is taken 17 GPa, and the Griffith energy
 320 release rate, μ_{g_c} , 100 Pa·m. For each value set of the statistical parameters, we simulate
 321 fracture propagation on 1,000 realizations of the random mechanical property fields

322 We include Fig. 4(a) for illustrative purposes, where we depict the contour plot of one
 323 realization of the Young modulus field of the inner domain with $C_v = 0.2$ and $l_c = 0.10$ m.
 324 We also plot the fracture path in white. For each parameter combination, we show the 1,000
 325 fracture trajectories Fig. 4(b) (gray color). In the same figure we also indicate the trajectory
 326 shown in Fig. 4(a) with a black solid line. With this raw data, we analyze the fracture path
 327 uncertainty in terms of C_v , and l_c . The mesh size is small enough to capture the spatial
 328 variability of the random fields imposed by l_c . In the following sections we describe our
 329 analyses and discuss our findings.

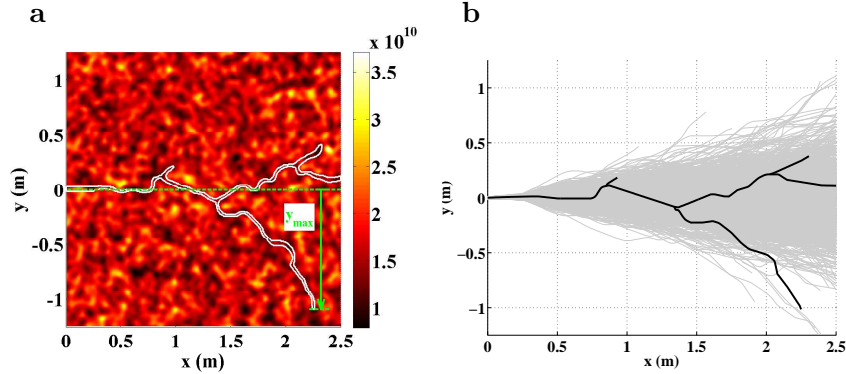


FIG. 4. Monte Carlo simulation framework. (a) We plot one realization of 2D log-normal Young modulus with a Gaussian correlation function. μ_E is 17 GPa, σ_E is 3.4 GPa, μ_{g_c} is 100 Pa·m, σ_{g_c} is 20 Pa·m, and l_c is 0.10 m. The fracture path is show in white. (b) Here we depict in gray the fracture paths of 1000 realizations of random E -fields, and in black the fracture path corresponding to the realization shown in panel (a).

TABLE I. Simulated data sets. The mean value of the Young modulus, μ_E , is taken 17 GPa, and the Griffith energy release rate, μ_{g_c} , 100 Pa·m.

Set	C_v	l_c (m)
1	0.1	0.10
2	0.2	0.10
3	0.3	0.10
4	0.2	0.05
5	0.2	0.15
6	0.2	0.20
7	0.2	0.40

330 IV.1. Influence of the coefficient of variation

331 We run 1,000 realizations for each considered value of C_v , with the correlation being $l_c =$
332 0.10 m. We study three values of $C_v = \{0.1, 0.2, 0.3\}$, resulting in three data sets of statistical
333 parameters of the mechanical properties. We simulate the fracture propagation for each
334 realization—a total of 3,000 paths. We will refer to y as the vertical distance between
335 the simulated fracture propagation and the deterministic one at a given x position. In
336 other words, the deviation of the fracture path from the straight, horizontal one. Since the

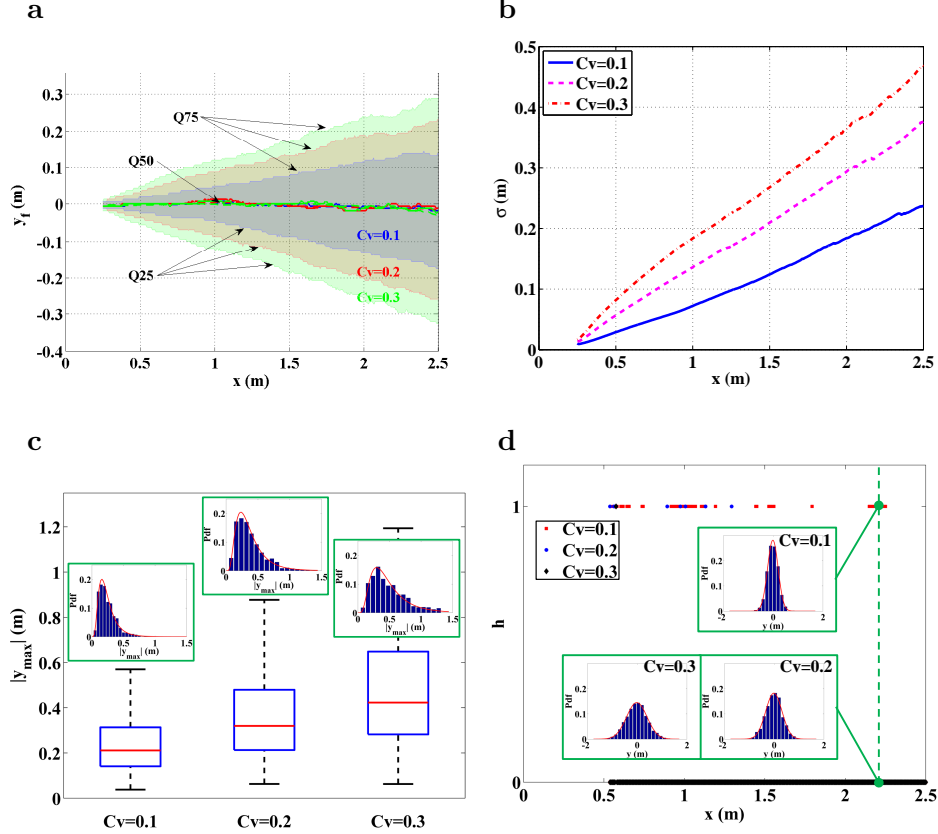


FIG. 5. Influence of C_v on the hydraulic fracture paths. (a) We plot the quartiles of the deviation, y , of the fracture path from the deterministic trajectory for three values of C_v . The initial "seed" fracture extends between points $(0, 0)$ and $(0.2, 0)$. A constant value of x , larger values of C_v implies a larger inter-quartile distance of y , *i.e.*, larger uncertainty. Also, for a given C_v , the inter-quartile distance increases linearly with the distance to the injection point. (b) Evolution of the standard deviation of y , σ , along the coordinate x , for three values of C_v . σ increases with the heterogeneity of the medium, and linearly with the distance to the injection point. (c) Box plot of the maximum deviation y_{max} of each fracture path for the three considered values of C_v . In the upper part of the box plots we show the histograms of y_{max} in blue and the fitted lognormal distributions in solid red line. The Kolmogorov-Smirnov test indicates that y_{max} can be modeled as lognormal in all cases. (d) We fit normal distributions to the deviation y at each x position, and plot the result of the Kolmogorov-Smirnov test to check whether the simulated data follow normal distributions—null hypothesis. We also depict three histograms in blue and the fitted normal distributions in solid red line.

TABLE II. Root-mean-square deviation, S_r , between the probability of the data sets 1 to 7 to the selected distributions: log-normal, Weibull, Gumbel and Gamma. The log-normal distribution provides the best fit followed by the gamma distribution.

Set	Log-normal	Weibull	Gumbel	Gamma
1	1.35×10^{-2}	3.24×10^{-2}	1.12×10^{-1}	1.99×10^{-2}
2	1.06×10^{-2}	3.44×10^{-2}	1.04×10^{-1}	2.18×10^{-2}
3	1.68×10^{-2}	2.86×10^{-2}	8.80×10^{-2}	1.93×10^{-2}
4	1.12×10^{-2}	3.82×10^{-2}	1.22×10^{-1}	2.29×10^{-2}
5	9.93×10^{-3}	3.31×10^{-2}	1.01×10^{-1}	2.99×10^{-2}
6	1.76×10^{-2}	3.83×10^{-2}	1.02×10^{-1}	2.99×10^{-2}
7	1.60×10^{-2}	2.51×10^{-2}	1.01×10^{-1}	1.65×10^{-2}

337 computational mesh in the propagation area is composed of 4-noded structured quadrilateral
 338 elements, there will be at least $(2.5 - 0.2)/0.008 = 288$ values of y for a single fracture path,
 339 where 2.5 m refers to the horizontal length of the propagation domain, 0.2 m is the length
 340 of the initial fracture, and 0.008 m is the mesh size.

341 Using all the simulated fracture paths for a given data set of statistical parameters, we
 342 compute the quartiles of the variable y at each x position of our computational mesh. We plot
 343 the quartiles for each value of C_v in Fig. 5(a), where the injection point is located at $(0, 0)$ and
 344 the initial fracture extends from that point to $(0.2, 0)$. The spatial variability —characterized
 345 by C_v — of the mechanical properties has a direct effect on the uncertainty of the fracture
 346 paths. For a fixed x position, larger values of C_v imply larger inter-quartile distance of y ,
 347 or in other words larger uncertainty. Moreover, for a given value of C_v , the inter-quartile
 348 distance linearly increases with the distance to the injection point. The second quartile
 349 —median— is roughly zero in all x positions, and the first and third quartiles are symmetrical
 350 with relation to the second one. These results indicate that fractures propagate with the
 351 same probability in the positive or negative parts of the y-axis, and that the mean behavior
 352 corresponds to the deterministic one.

353 The standard deviation of y, σ , at a given x position increases with C_v (Fig. 5(b)), which
 354 is also reflected by the evolution of the second and third quartiles with the x coordinate
 355 (Fig. 5(a)). Moreover, for a given value of C_v , the growth of the variance with the x-
 356 coordinate is linear. The mean of the deviation—not shown in the figure—is zero in all

357 vertical sections.

358 We determine the maximum absolute value of the deviation, y_{max} , of each fracture path.
359 We depict the box plots of y_{max} for the three considered values of C_v in Fig. 5(c). The
360 spatial variability has also a direct influence on the extremal values of y : the median, the
361 inter quartile range and the whisker length increase with C_v . The evolution of the quartiles
362 is rather linear with C_v , in the sense that the variability of y_{max} increases with C_v . The
363 upper part of the boxes is bigger than the lower one, which indicates that the probability
364 distributions of y_{max} are asymmetric.

365 We fit log-normal, Weibull, Gumbell and gamma probability distribution functions
366 (PDFs) to the three data sets of y_{max} . We assess the adjustment through the root-mean-
367 square deviation, S_r , between the probability of the data to the selected distribution, see
368 Table II. The results indicate that the log-normal PDF provides the best fit to the three
369 data sets.

370 We plot the histograms of y_{max} in the insets of Fig. 5(c) and the fitted lognormal PDFs
371 to the data in red solid lines. We perform a Kolmogorov-Smirnov test to check whether the
372 simulated data come from log-normal distributions –null hypothesis–. The test indicates
373 that the null hypotheses are not rejected at the 5% of significance level in any set. The
374 histograms also show that the mean and the variance of y_{max} increase with C_v .

375 We analyze the probability distribution of the deviation at each x position. We fit
376 normal distributions to the three data sets at each vertical cross section, and we perform the
377 Kolmogorov-Smirnov test to check whether simulated data follow normal distributions—null
378 hypothesis. The results of the test (Fig. 5(d)) show that in most cases the null hypotheses
379 cannot be rejected at the 5% of significance level. Moreover, the number of rejected cases
380 decrease as the value of C_v increases. Nevertheless, the rejected cases qualitatively may
381 follow normal distributions. We show this by plotting in blue the histograms of y at the
382 section $x = 2.25$ m for the three considered values of C_v , and in solid red line the fitted normal
383 distributions. The test rejects the set with $C_v = 0.1$, although these data qualitatively seem
384 to be normally distributed.

385 We have found that the deviation of the fracture path from the deterministic trajectory
386 increases with the heterogeneity of the medium. Moreover, at every position of the deter-
387 ministic fracture, the position of the random trajectories follows a normal distribution whose
388 mean is the deterministic path and the standard deviation increases with the distance to
389 the origin of the fracture. Finally, we have found that this standard deviation linearly in-

390 creases with the distance, which suggests that fracture trajectories can be described through
391 random walks.

392 **IV.2. Influence of the correlation length**

393 We investigate the uncertainty of the fracture propagation occasioned by the correlation
394 length, l_c , in five data sets of statistical parameters with $l_c = \{0.05, 0.10, 0.15, 0.20, 0.40\}$ m
395 whereas C_v is constant and equals to 0.2. We run 1000 realizations for each data set.

396 We plot the quartiles of y at each x position of the computational mesh for three smallest
397 values of l_c in Fig. 6(a). The influence of the correlation length on fracture trajectories
398 is qualitatively similar to that of C_v : for a fixed position, higher values of l_c imply larger
399 inter-quartile distance, or in other words, larger uncertainty in the fracture path. As in
400 the previous analysis, the inter-quartile distance linearly increases with the distance to the
401 injection point. In the same manner, the second quartile –median– is approximately zero in
402 all sections, and the first and third quartiles are also symmetric with respect to the median.

403 We also compute the quartiles for larger values of l_c , 0.20 and 0.40 m,—not shown in
404 Fig. 6(a)—, but their values are quite similar to those of $l_c = 0.15$ m. We attribute this
405 behavior to the ratio between l_c and the size of the propagation domain. The correlation
406 length is a way to consider the size of the inclusions in the rock –rigid or soft–, and the
407 coefficient of variation accounts for the variability in the magnitude of the mechanical prop-
408 erties between the soft and rigid inclusions. As a result, the size of the domain must contain
409 an enough number of inclusions to properly capture the variability. This number can be
410 estimated as the ratio between the area of the domain and the mean area of one inclusion,
411 considered these ones as circles with radius $\sqrt{2}l_c$ according to Eq. (3). In our square domain
412 of side 2.5 m and $l_c = \{0.05, 0.10, 0.15, 0.20, 0.40\}$ m, the estimated number of inclusions
413 is respectively $\{398, 99, 44, 25, 6\}$. For values of l_c larger than 0.15 m, the domain does not
414 seem to be big enough to capture the spatial variability of the mechanical properties, since
415 the effect of l_c over the deviation of the fractures does not increase. For even larger val-
416 ues of l_c , the domain could become part of just one inclusion and no variability would be
417 captured, *i.e.*, the problem would become deterministic.

418 The standard deviation of y at a given section increases with l_c (Fig. 6(b)). Moreover, for
419 a given value of l_c the standard deviation linearly increases with the distance to the injection
420 point. The mean of the deviation –not shown in the figure– is zero at all vertical sections.

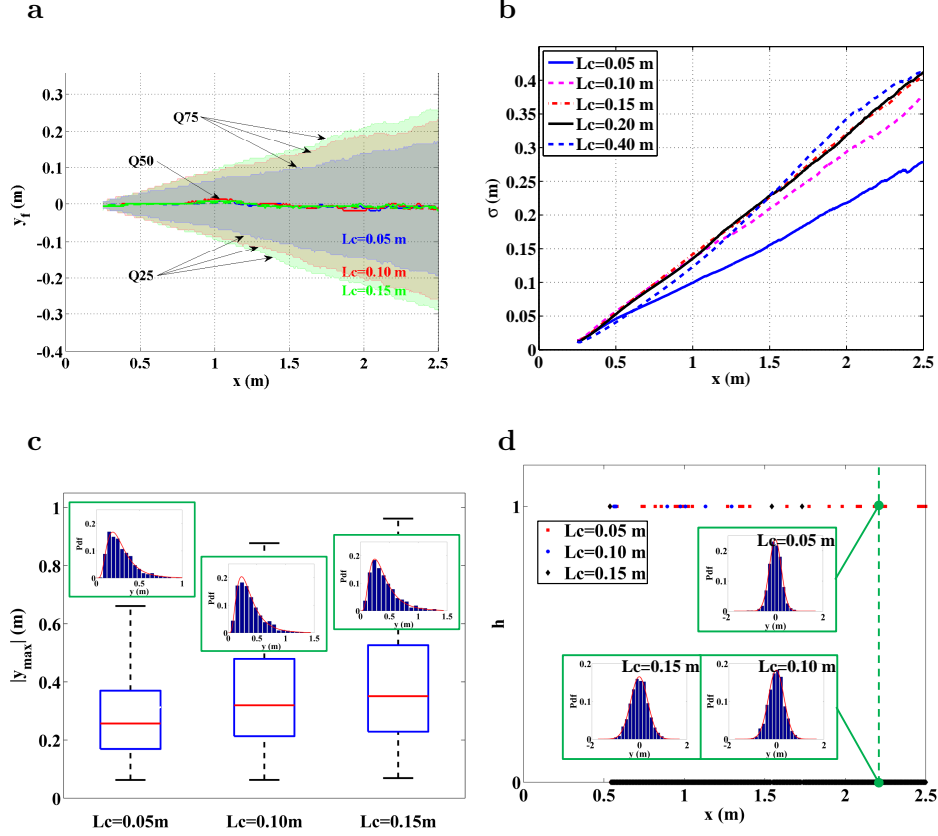


FIG. 6. Influence of l_c on the hydraulic fracture path. (a) Quartiles of the deviation y of fracture paths from the deterministic trajectory for three values of the correlation length, l_c . For a constant value of x , larger values of l_c lead to larger inter-quartile distance of y , *i.e.*, more uncertainty. Also, for a given l_c , the inter-quartile distance increases linearly with distance to the injection point. (b) Evolution of the standard deviation of y — σ —with the coordinate x for five values of l_c . For a given x , σ increases with l_c up to $l_c = 0.15$ m, and then σ remains constant for larger values. Also, σ increases with the distance to the injection point, suggesting that fractures deviate further from the straight path as either l_c or the distance to the injection point increase. (c) Box plot of the maximum deviation y_{max} of each fracture path for the three considered values of l_c . In the upper insets we show the histograms of y_{max} in blue and the fitted log-normal distributions in solid red line. The Kolmogorov-Smirnov test indicates that y_{max} can be modeled with lognormal distributions in all cases. (d) Here we plot the result of the Kolmogorov-Smirnov test to check whether simulated deviations y at each x position follow normal distributions—null hypothesis. We also show three histograms in blue and the fitted normal distributions in solid red line.

421 We also plot the evolution of σ for $l_c = 0.20$ m and 0.40 m, whose aspect is essentially
422 identical to the set with $l_c = 0.15$ m

423 We compute y_{max} of each fracture path and depict the box plots of y_{max} for the three
424 values of l_c in Fig. 6(c). The median, the inter quartile range and the whisker length increase
425 with l_c , but the rate of the growth decreases for the largest values of l_c . The probability
426 distributions of y_{max} are asymmetric, and the asymmetry seems to increase with l_c . We fit
427 log-normal, Weibull, Gumbell and gamma PDFs to the three data sets of y_{max} and assess
428 the adjustments through the root-mean-square deviation, Table II. The results indicate that
429 the log-normal PDF provides the best fit to the three data sets. We plot the histograms
430 of y_{max} in the insets of Fig. 6(c) and the fitted log-normal PDFs to the data in solid red lines.
431 The Kolmogorov-Smirnov test indicates that the null hypotheses—data follow log-normal
432 distributions—are not rejected at the 5% of significance level in any set.

433 We study the probability distribution of y at each x position. We fit normal distribu-
434 tions to the three data sets at each vertical cross section. The results of the Kolmogorov-
435 Smirnov test (Fig. 6(d)) show that in most cases the null hypothesis –data follow normal
436 distributions– cannot be rejected at the 5% of significance level. As in the previous analysis,
437 the number of rejected cases decrease as the value of l_c rises. Nevertheless, the rejected
438 cases qualitatively may follow normal distributions: we plot in blue the histograms of y at
439 the section $x = 2.25$ m for the three values of l_c , and in solid red line the adjusted normal
440 distribution. The test rejects the set of $l_d = 0.05$ m, although apparently follows a normal
441 distribution.

442 Our main finding is that the correlation length has a similar effect as the heterogeneity
443 of the medium. However, its impact increase with its magnitude up to a value from which
444 remains constant. We have attributed this effect to the ratio between l_c and the size of the
445 domain, for large values of l_c the domain is not big enough to capture the variability.

446 V. CONCLUSIONS

447 A fluid-driven fracture propagating in an homogeneous and isotropic impermeable elastic
448 medium follows a straight trajectory. In contrast, real materials have heterogeneous me-
449 chanical properties which induce deviations of the fracture trajectories. We have analyzed
450 the influence of the spatial variability of mechanical properties on the complexity and spa-
451 tial structure of the trajectories of fluid-driven fractures. In particular, we characterized the

452 deviations of fracture trajectories from the straight, deterministic path. For this purpose,
453 we have generated lognormal isotropic random fields of the Young modulus and the Griffith
454 energy release rate. These random correlated fields are statically defined by the coefficient
455 of variation, C_v —ratio between the standard deviation and the mean of the field—and the
456 correlation length, l_c —a characteristic length which defines the spatial correlation of the
457 fields.

458 To focus on the crack growth process, the geometry of our case study is a square domain
459 of side 2.5 m, where an initial horizontal fracture of 0.2 m length is located at the center
460 of the left side. To avoid boundary effects, this refined inner domain is embedded inside a
461 larger square domain of 60 m. We simulate the pressure-driven propagation of the initial
462 fracture inside the inner domain by conducting numerical simulations based on the phase-
463 field approach under plane strain conditions. Based on these simulations, we studied the
464 deviation of the trajectory of each fracture, $y(x)$ from the deterministic path at several
465 points along the trajectory.

466 We considered three values of the coefficient of variation, C_v , with constant l_c , running
467 a thousand realizations of the random mechanical property fields for each value of C_v , and
468 simulated the fracture propagation for each realization. For a fixed position located along the
469 horizontal deterministic fracture path, x , the mean and median of y at that point are equal
470 to the deterministic value, and the first and third quartiles are symmetrical in relation to the
471 median. However, the standard deviation of y at a given position increases with C_v , and for
472 a given C_v it increases with the distance to the initial fracture. We have also analyzed the
473 probability distribution of y at several positions located in the deterministic fracture path,
474 and we have found that in most of them y can be well represented by normal distributions.
475 We have also found that the absolute maximum value of y of each fracture trajectory follows
476 log-normal distributions whose mean and standard deviation increase with C_v .

477 We have conducted several analyses to estimate the influence of the correlation length l_c
478 on the fracture trajectories. We have considered five values of l_c , with constant C_v , and
479 run a thousand realizations for each value. We have found that the smallest values of l_c
480 have a similar effect as C_v , *i.e.*, the variability of the fracture trajectories increases with l_c .
481 However, the largest values of l_c have resulted into paths with similar variability. We have
482 attributed this behavior to the ratio between the size of the domain and l_c . Since the domain
483 has constant size, it is unable to properly capture the spatial variability of the mechanical
484 properties for the largest values of l_c . Moreover, for values of l_c comparable to the size of

485 the domain, the fields can even become deterministic.

486 Our research shows that heterogeneity in the mechanical properties of fracturing materials
487 significantly influences the complexity of fluid-driven fracture trajectories. Not surprisingly,
488 the deviation from the deterministic paths predicted by homogeneous media theories in-
489 creases with the variability of mechanical properties. This is important in practice for the
490 application of hydraulic fracturing to the stimulation of geothermal reservoirs or the recov-
491 ery of oil and gas from unconventional reservoirs. In the latter, one of the major concerns is
492 the potential for groundwater contamination due to the unintentional hydraulic connection
493 of the target reservoir with surrounding fresh water bodies [11, 12]. These hydraulic con-
494 nections may arise from stimulated hydraulic fractures that deviate significantly from their
495 expected fracture paths due to heterogeneity. Our research can be the starting point for fur-
496 ther developments of statistical procedures to estimate the likelihood that a fracture reaches
497 an aquifer during the hydraulic fracturing processes, among other applications. Moreover,
498 the effect of anisotropy in the random mechanical properties or in the initial stress field are
499 also important open questions.

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