

1 **ASSESSING FUTURE GREENHOUSE GAS CONCENTRATIONS IN MADRID,**  
2 **SPAIN: A DYNAMICAL DOWNSCALING APPROACH USING WRF-VPRM**

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7 Preprint: Air Quality, Atmosphere & Health (2025) 18:3813–382 .[https://doi.org/10.1007/s11869-025-](https://doi.org/10.1007/s11869-025-01798-1)  
8 [01798-1](https://doi.org/10.1007/s11869-025-01798-1)cities

9 **ACKNOWLEDGEMENTS**

10 DISTENDER has received funding from the European Union’s Horizon EU research and innovation  
11 programme under grant agreement No 101056836.

12 The UPM co-authors acknowledge the computer resources and technical assistance provided by  
13 the Centro de Supercomputación y Visualización de Madrid (CeSViMa). The UPM co-authors  
14 thankfully acknowledge the computer resources, technical expertise and assistance provided by  
15 the Red Española de Supercomputación (RES)

16  
17  
18 **ABSTRACT**

19 This study employs a comprehensive modeling approach that combines the Weather Research Forecast  
20 (WRF) model with the Vegetation Photosynthesis and Respiration Model (VPRM) to gain insights into the  
21 intricate dynamics of GHGs flux and concentration. The primary objective is to conduct high-resolution  
22 simulations of GHGs transport, dispersion, and concentrations spanning from 2015 to 2050 under future  
23 climate scenario over Madrid (Spain) region. The global climate scenarios datasets are provided by the 6th  
24 Coupled Model Intercomparison Project (CMIP6). Dynamical downscaling is done by the WRF model to  
25 increase the spatial resolution of the climate data. The results of the atmospheric simulations illuminate the  
26 impact (future – present) of emission source distribution and climate data on GHG concentrations.  
27 Additionally, the incorporation of the vegetation photosynthesis and respiration model provides valuable  
28 perspectives on the role of terrestrial ecosystems in the global carbon cycle and their influence on  
29 atmospheric CO<sub>2</sub> concentrations.

30  
31 **Keywords:** Climate Change, CO<sub>2</sub>, VPRM  
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## 1. INTRODUCTION

Carbon dioxide (CO<sub>2</sub>) is a long-lived greenhouse gas (GHG) in the atmosphere that plays an important role in the Earth's climate system. Concentrations of CO<sub>2</sub> in the atmosphere come from both anthropogenic and natural sources, as well as anthropogenic (fuel combustion) and natural sources (biogenic fluxes from ecosystems). In recent decades, atmospheric concentrations of CO<sub>2</sub> have increased significantly, largely due to human activity as anthropogenic CO<sub>2</sub> emissions exceed the CO<sub>2</sub> sequestration capacity of terrestrial ecosystems and oceans. Quantifying the exchange of CO<sub>2</sub> between the atmosphere and terrestrial ecosystems (biogenic fluxes) is essential for understanding the carbon balance and accurately identifying sources and sinks. Urban areas, in particular, are the largest contributors to global CO<sub>2</sub> emissions, accounting for more than 70% of total anthropogenic emissions (IPCC,2014). In response, many cities around the world have committed to achieving net GHG emissions by 2050. However, achieving this target requires robust methodologies for the quantification of anthropogenic (Davis et al., 2017) and biogenic CO<sub>2</sub> emissions with high spatial resolution. From this information, effective mitigation actions can be designed. Anthropogenic emission inventories have been developed to estimate CO<sub>2</sub> emissions in all sectors, including industry, transport, residential and commercial activities ( Sówka and Bezyk, 2018). These inventories are continuously refined to improve temporal and spatial resolution. In parallel, natural fluxes - especially those driven by interactions between the biosphere and the atmosphere - introduce additional complexity into CO<sub>2</sub> concentration dynamics. These biogenic fluxes are highly sensitive to meteorological conditions, such as temperature, solar radiation and land cover type, and can significantly influence CO<sub>2</sub> balances. Despite advances in observation networks and modelling capabilities, CO<sub>2</sub> fluxes at regional/local scales remain insufficiently characterised due to their large spatial and temporal variability, especially in heterogeneous environments with complex urban and vegetation mosaics. To address this challenge, mesoscale atmospheric models such as the Weather Research and Prediction coupled with Chemistry and the Plant Respiration and Photosynthesis Model (WRF/Chem-VPRM) have been developed. These models allow simultaneous simulation of meteorological fields, CO<sub>2</sub> transport and biospheric exchange with high spatial and temporal resolution.

While global climate models (GCMs), such as those used in the Sixth Assessment Report of the Intergovernmental Panel on Climate Change (IPCC AR6), provide valuable information on global trends under various Shared Socio-Economic Pathways (SSPs), their coarse resolution limits their usefulness for analysing local and regional climate impacts. GCMs often do not capture fine-scale phenomena, such as local wind patterns, land-use heterogeneity or topographic effects, which are crucial for accurate CO<sub>2</sub> modelling. Downscaling techniques, in particular dynamic downscaling with regional climate models (RCMs), are therefore essential for translating global climate projections into practical information at regional scales. Dynamic downscaling allows RCMs to incorporate regional physiographic features such as elevation, vegetation and urban structure, improving their ability to simulate climate variables and their corresponding effects on biogenic emissions and CO<sub>2</sub> concentrations (Dong et al., 2021). In this context, the present study investigates the high-resolution impact of future climate scenarios on CO<sub>2</sub> fluxes in the Iberian Peninsula, Community of Madrid and Municipality of Madrid, using a nested modelling approach with the WRF/Chem-VPRM model driven by the climate scenarios that are dynamically rescaled, offering a multi-scale perspective of carbon dynamics. The analysis covers the period from 2015 to 2050 and focuses on isolating the influence of future climate conditions on biogenic CO<sub>2</sub> emissions and total atmospheric CO<sub>2</sub> concentrations. To ensure consistency, anthropogenic emissions and land use are held constant at 2018 levels, allowing the climate signal to emerge as the dominant driver of change. This methodology offers a unique opportunity to assess how future climate changes - under different SSPs - may modulate biospheric CO<sub>2</sub> exchange and impact CO<sub>2</sub> mitigation strategies.

Although CMIP6 (Eyring et al., 2016) GCMs represent a major step forward in climate modelling, offering better spatial resolution and more sophisticated physical parameterisations compared to their CMIP5 predecessors, they still exhibit significant biases in regional climate simulation, especially in complex terrain or extreme conditions ( Zhu and Yang, 2020; Hu and Zhou, 2021; Cui et al., 2021). These limitations underline the importance of regional downscaling in producing reliable climate information for urban- and regional-scale carbon modelling. In this study, the use of rescaled CMIP6 climate scenarios in combination with high-resolution CO<sub>2</sub> modelling addresses a critical knowledge gap: how future climate change might influence biospheric CO<sub>2</sub> emissions in an urban context. By assessing climate-driven changes in net carbon exchange in the Madrid region, this work contributes to a better understanding of the interactions between regional climate dynamics and the carbon cycle, an essential step towards designing effective mitigation strategies for climate-resilient cities.

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101 In this study, four climate scenarios from CMIP6 were considered to assess the potential impacts of future  
102 climate conditions on biogenic CO<sub>2</sub> emissions and concentrations. These scenarios correspond to the Tier  
103 1 Shared Socioeconomic Pathways (SSPs), which combine different trajectories of socioeconomic  
104 development (Meinshausen et al., 2020) with associated greenhouse gas emissions and radiative forcing  
105 levels projected for the year 2100. The selected scenarios—SSP1-2.6, SSP2-4.5, SSP3-7.0, and SSP5-8.5—  
106 span a wide range of plausible future climate outcomes and represent different levels of mitigation and  
107 adaptation challenges (O'Neill et al., 2016; Riahi et al., 2017).

108

109 SSP1-2.6 represents a sustainability-oriented pathway, often referred to as the "green road." It assumes low  
110 challenges to both mitigation and adaptation, with strong international cooperation, rapid technological  
111 advancement, and a shift towards environmentally friendly energy systems. The scenario targets a low  
112 radiative forcing of 2.6 W/m<sup>2</sup> by 2100, roughly consistent with the Paris Agreement goal of limiting global  
113 warming to below 2°C. It is analogous to the RCP2.6 pathway used in CMIP5 and reflects ambitious climate  
114 protection efforts.

115

116 SSP2-4.5, known as the "middle-of-the-road" scenario, assumes intermediate challenges to mitigation and  
117 adaptation. It envisions a world that follows historical patterns, with moderate progress in sustainability  
118 and climate policy. This pathway results in a radiative forcing of 4.5 W/m<sup>2</sup> by the end of the century and  
119 serves as a baseline scenario for many climate impact assessments. It is considered a direct update to  
120 RCP4.5 from the previous CMIP5 framework.

121

122 SSP3-7.0 portrays a fragmented world characterized by regional rivalry and slow economic development.  
123 Under this pathway, national priorities dominate over global cooperation, and mitigation efforts are limited.  
124 As a result, greenhouse gas emissions rise substantially, leading to a radiative forcing of 7.0 W/m<sup>2</sup> by 2100.  
125 This scenario reflects a high-challenge context for both mitigation and adaptation and is indicative of policy  
126 failure in achieving international climate targets.

127

128 SSP5-8.5 represents a fossil-fueled development pathway, often described as the "highway" scenario. It  
129 envisions rapid economic growth powered by continued reliance on carbon-intensive energy sources such  
130 as coal, oil, and natural gas. While adaptation capacity may increase due to technological advancement and  
131 economic strength, mitigation efforts remain minimal. This scenario leads to a radiative forcing of 8.5 W/m<sup>2</sup>  
132 by 2100, making it the most extreme of the four in terms of climate change impacts and comparable to  
133 RCP8.5 in CMIP5. It serves as an upper-bound scenario for assessing worst-case outcomes (Hausfather &  
134 Peters, 2020).

135

136 Previous studies have demonstrated the value of WRF-VPRM modelling system in capturing biogenic CO<sub>2</sub>  
137 dynamics (Mahadevan et al., 2008; Ahmadov et al., 2009). This model is based on satellite-derived  
138 vegetation indices and meteorological data obtained by WRF to simulate two key components of biogenic  
139 carbon flux: gross ecosystem exchange (GEE) and ecosystem respiration (ER). Numerous applications of  
140 WRF/Chem-VPRM in Europe, Asia and North America have demonstrated their ability to capture  
141 mesoscale transport patterns and vegetation-driven variability (Ballav et al., 2012; Pillai et al., 2016; Park  
142 et al., 2018). Urban ecosystems, in particular, exhibit strong seasonal and diurnal variability in CO<sub>2</sub> uptake,  
143 influenced by vegetation phenology (Hutyra et al., 2014), microclimate and anthropogenic activity (;  
144 Hardiman et al., 2017; Miller et al., 2020; Bezyk et al., 2021). Consequently, future changes in temperature,  
145 radiation and moisture availability may significantly alter the balance between biogenic uptake and release,  
146 especially during extreme events such as heat waves or droughts.

147

## 148 2. METHODOLOGY

149

### 150 2.1 Experiment

151 To assess the impact of future climate conditions on biogenic emissions and atmospheric CO<sub>2</sub> dynamics, a  
152 comprehensive numerical experiment was carried out using the WRF/Chem-VPRM modelling system,  
153 which will be described in more detail in the next section. The simulations cover a nested domain  
154 configuration centred (40.478°N, 3.704°W ) over Madrid (Spain) covering the Iberian Peninsula, the  
155 Community of Madrid and the Madrid metropolitan area. The model employs a two-way interactive nesting  
156 approach across three spatial domains: Domain 1 (D1): 27 km horizontal resolution. Domain 2 (D2): 9 km

horizontal resolution Domain 3 (D3): 3 km horizontal resolution Each domain is structured as a 30 × 30 grid with 33 vertical atmospheric layers extending from the surface to 10 hPa.

The simulation period spans 35 years, from 2015 to 2050, with a start-up period of one day used to initialise the atmospheric state each year. The simulations were run continuously for one year, generating hourly results for each of the three spatial domains and the four CMIP6 Tier 1 climate scenarios (SSP1-2.6, SSP2-4.5, SSP3-7.0 and SSP5-8.5). The computational workload was distributed over 120 parallel jobs, each running on 9 cores, accumulating a total of approximately 50,000 CPU hours and producing over 10 TB of output data.

The climate forcing for the WRF/Chem-VPRM simulations was obtained from the six-hourly outputs available in the CMIP6 repositories of the Max Planck Institute Earth System Model (MPI-ESM1.2-HR), an Earth system model that has an atmospheric component based on ECHAM6.3, with a horizontal resolution of T127 (approximately 0.93° × 0.93° or ~103 km) and 95 vertical levels extending down to 0.01 hPa. The ocean model is MPIOM1.63, with a resolution of TP04 (0.4° × 0.4° or ~44 km) and 40 vertical levels. This particular GCM was selected because it is explicitly tuned to a 3 K equilibrium climate sensitivity, and a well-balanced radiation budget (Müller et al., 2018). These features make MPI-ESM1.2-HR particularly suitable for regional downscaling and impact assessment applications. Historical climate simulations cover the period 1980-2014, while future projections cover 2015-2100. For this study, the period from 2015 to 2050 was extracted and dynamically rescaled, and initial and lateral boundary conditions for meteorological variables were updated every six hours using the output of MPI-ESM1.2-HR.

## 2.2 Modelling system

The simulation chain begins with the conversion of CMIP6 six hours GCM outputs (2015-2050) into WRF-compatible intermediate files using the 'cmip6-to-wrffinterm' Python tool (<https://github.com/lzhenn/cmip6-to-wrffinterm>), which will be read by the WRF Preprocessing System (WPS) with the 2018 (reference year) land use and topographic data. Next, WRF/Chem-VPRM model is executed using the initial and boundary conditions from WPS and 2018 anthropogenic emissions. During the simulation (on-line) VPRM calculates biogenic CO<sub>2</sub> fluxes at each time step using vegetation indices, simulated temperature, and radiation. These fluxes are then transported by the chemistry module within WRF/Chem. WRF/Chem is a fully compressible, non-hydrostatic meteorological (WRF) and chemical transport (Chem) model developed by NCAR (Grell et al., 2005). It integrates atmospheric dynamics and chemical processes in an online framework, ensuring consistency in physics, transport, and emissions across the meteorological and chemical components (Powers et al., 2017). In this study, WRF-Chem version 4.1.5 is used, with the following physical parameterization: Microphysics: Morrison double-moment scheme (Morrison et al., 2009). Radiation: Rapid Radiative Transfer Model for GCMs (RRTMG) for both shortwave and longwave (Iacono et al. 2008). Planetary Boundary Layer (PBL): Yonsei University (YSU) scheme (Hong et al. 2006). Surface Layer: Monin-Obukhov scheme (Monin and Obukhov 1954). Land Surface: Noah LSM (Tewari et al., 2004). Cumulus Parameterization: Grell 3D Ensemble (G3) (Grell and Dévényi 2002). CO<sub>2</sub> Chemistry: Passive tracer mode (no chemical reactions).

Anthropogenic CO<sub>2</sub> emissions are based on the CAMS-GLOB-ANT emission inventory, downscaled using the EMIMO emission model developed at UPM (San Jose et al, 2015). This inventory provides hourly gridded emissions at 0.1° × 0.1° resolution for the year 2018. The biogenic emissions are calculated by VPRM module. The VPRM (Mahadevan et al., 2008; Ahmadov et al., 2007) is a model that estimates Net Ecosystem Exchange (NEE) of CO<sub>2</sub> as the difference between Gross primary production (GPP, uptake) and ecosystem respiration (R, release). GPP and R are calculated by the equations 1 and 2.

$$GPP = \lambda \times T_{scale} \times P_{scale} \times W_{scale} \times FAPAR_{PAV} \times [PAR / (1 + PAR / PAR_0)] \quad (1)$$

$$R = \alpha \times T_{air} + \beta \quad (2)$$

Where:

$\lambda$ ,  $\alpha$ ,  $\beta$ ,  $PAR_0$  are empirical parameters dependent on the eight land cover types, optimized by Europe.

$T_{air}$  is 2 m air temperature from WRF

PAR is the Photosynthetically Active Radiation, derived as 0.5 × downward shortwave radiation (SW) from WRF.

$T_{scale}$ ,  $P_{scale}$  and  $W_{scale}$  are three dimensionless scaling factors related with the temperature sensitivity of photosynthesis, effects of leaf age on canopy photosynthesis and water stress effect respectively. They are dependents from EVI (Enhanced Vegetation Index) and LSWI indices (Land Surface Water Index).

217 These indices can be derivated from the reflectance data from the MODIS-Terra MOD09A1 v6 product,  
 218 with 8-day temporal resolution and 500 m spatial resolution. The applied equations are (3) for EVI and (4)  
 219 for LSWI:

$$220 \quad \text{EVI} = 2.5 \times (\rho_{\text{NIR}} - \rho_{\text{RED}}) / (\rho_{\text{NIR}} + 6 \times \rho_{\text{RED}} - 7.5 \times \rho_{\text{BLUE}} + 1) \quad (3)$$

$$222 \quad \text{LSWI} = (\rho_{\text{NIR}} - \rho_{\text{SWIR}}) / (\rho_{\text{NIR}} + \rho_{\text{SWIR}}) \quad (4)$$

223  
 224 Where  $\rho$  values represent surface reflectance in the blue, red, near-infrared (NIR), and shortwave infrared  
 225 (SWIR) bands.

### 227 3. RESULTS

228  
 229 Before presenting the results of the CO<sub>2</sub>-related simulations, an evaluation of the modelling system has  
 230 been carried out using the temperature variable, due to its importance and data availability.

#### 232 3.1 Evaluation

233 To evaluate the performance of the WRF/Chem-VPRM model in simulating near-surface air temperature  
 234 (2 meters high) , we used hourly observed temperature data from three meteorological stations located  
 235 within Domain 3 (3 km resolution). The stations are identified by: Barajas, Cuatro Vientos and Getafe  
 236 which are spatially distributed across the domain. Hourly data were available from 2015 to 2024, a total of  
 237 72,959 valid hourly observations per station. The hourly values of the 3 stations (station 0) were also  
 238 averaged to be compared with the mean values of the three cells where the stations are located. For each  
 239 model simulation corresponding to the four CMIP6-based climate scenarios (SSP1-2.6, SSP2-4.5, SSP3-  
 240 7.0, and SSP5-8.5), model results were compared with observed data. The comparison included two  
 241 standard performance metrics, such as the coefficient of determination ( $R^2$ ), and the normalised mean bias  
 242 (NMB). The results are showed in Table 1, the table below summarizes the NMB and  $R^2$  values for each  
 243 scenario and station  
 244

245 Table 1: Performance metrics

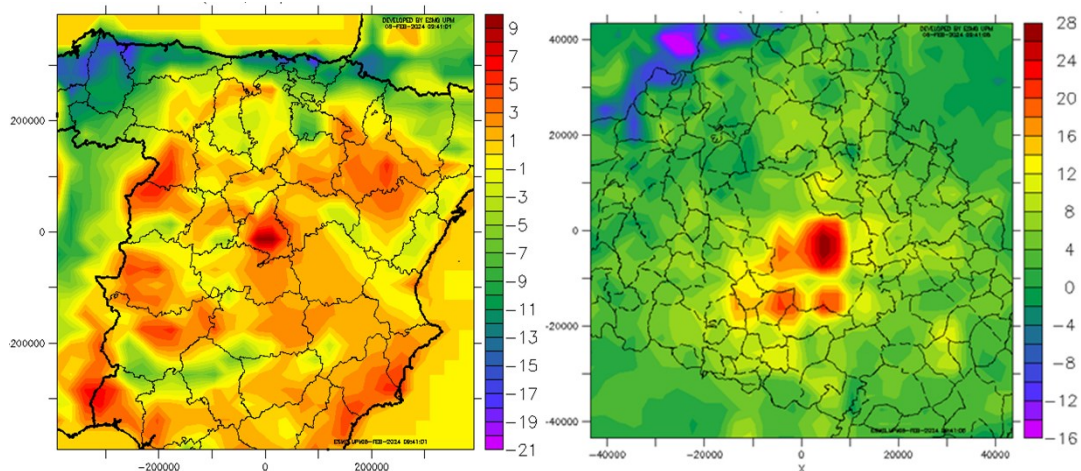
Simulation	Station	NMB	$R^2$
<b>ssp126</b>	0	0,44	0,84
<b>ssp245</b>	0	0,2	0,82
<b>ssp370</b>	0	0,41	0,83
<b>ssp585</b>	0	0,36	0,84
<b>ssp126</b>	1	0,33	0,84
<b>ssp245</b>	1	0,43	0,83
<b>ssp370</b>	1	-0,2	0,83
<b>ssp585</b>	1	0,17	0,85
<b>ssp126</b>	2	0,24	0,81
<b>ssp245</b>	2	0,1	0,81
<b>ssp370</b>	2	-0,2	0,81
<b>ssp585</b>	2	0,04	0,82
<b>ssp126</b>	3	0,32	0,84
<b>ssp245</b>	3	0,16	0,83
<b>ssp370</b>	3	-0,4	0,84
<b>ssp585</b>	3	0,17	0,85

249 Overall, the WRF/Chem-VPRM model shows a high ability to reproduce the temporal evolution of hourly  
250 temperature, with  $R^2$  values above 0.80 for all scenarios and stations. These high correlation coefficients  
251 indicate that the model is able to accurately capture the variability and trends in observed temperature over  
252 the eight-year period. In particular, the best agreement was found for the SSP5-8.5 scenario, which achieved  
253 the highest  $R^2$  values (up to 0.85) and the lowest bias values, suggesting that it is the most representative of  
254 recent climate conditions in the region. We can therefore consider SSP5-8.5 as the reference scenario or  
255 Business As Usual (BAU).

256  
257 However, despite the high correlation, a consistent positive bias is observed in most simulations, where  
258 modelled temperatures generally underestimate actual values. This bias is present in all scenarios except  
259 SSP3-7.0, where a negative bias is recorded. It is important to note that the grid-cell-based nature of the  
260 model introduces an inherent spatial averaging, which can dampen the magnitude of extreme events  
261 compared to point-based observations. Increasing the spatial resolution above 3 km could improve the  
262 model's ability to capture these more point-based variations, however, given the 35-year simulation horizon  
263 and associated computational burden this refinement is beyond the scope of the present work.

### 266 3.2 Spatial distribution

267 The following images include a set of maps showing the results derived from computational simulations.  
268 These images clarify the spatial dispersion patterns of various simulated parameters across various spatial  
269 scales. Each map provides insight into the spatial distribution of the simulated variables, outlining their  
270 respective patterns and variations across the study area. Figure 1 shows the spatial distribution of the CO<sub>2</sub>  
271 emission budget, comprising the cumulative sum of biogenic (which may be negative in some areas in case  
272 of absorption dominance) and anthropogenic emissions, in the context of the SSP5-8.5 climate scenario.  
273 The figure corresponds to simulations performed at spatial resolutions of 27 km (left) and 3 km (right),  
274 respectively. The data shown are mean values derived from hourly data averaged over the entire 35-year  
275 simulation period.



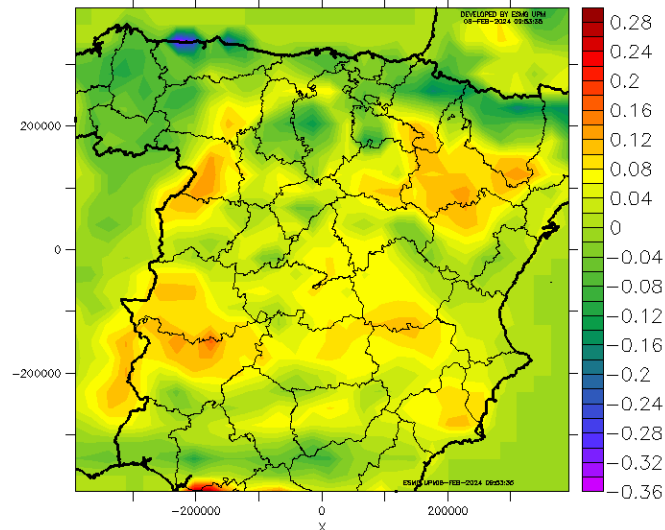
276 **Figure 1.** Spatial distributions of simulated CO<sub>2</sub> emissions (kmol km<sup>-2</sup> hr<sup>-1</sup>) budget (biogenic +  
277 anthropogenic contributions). Mean of 35 years 2015-2050 from WRF/Chem-VPRM simulation (27 km)  
278 for SSP585 scenario.  
279

280 In Figure 1, the Iberian Peninsula (left) shows a combination of areas with a negative net balance, indicative  
281 of CO<sub>2</sub> absorption, together with areas characterised by predominant CO<sub>2</sub> emissions, as clearly seen in the  
282 central area of the image where the city of Madrid is located, where anthropogenic emissions are very high.  
283 In addition, the northern regions of Spain and Portugal present important CO<sub>2</sub> sinks as they are areas of  
284 dense vegetation. Right part of Figure 1 provides a more detailed examination of the Madrid region,  
285 revealing a predominance of positive net balances, particularly important in the urban area of Madrid. In  
286 contrast, only a few isolated sites serve as CO<sub>2</sub> sinks, albeit with lower magnitudes than in the north of the  
287 Peninsula. These sinks are located in the mountainous area of Madrid, where the density of vegetation is  
288 higher.

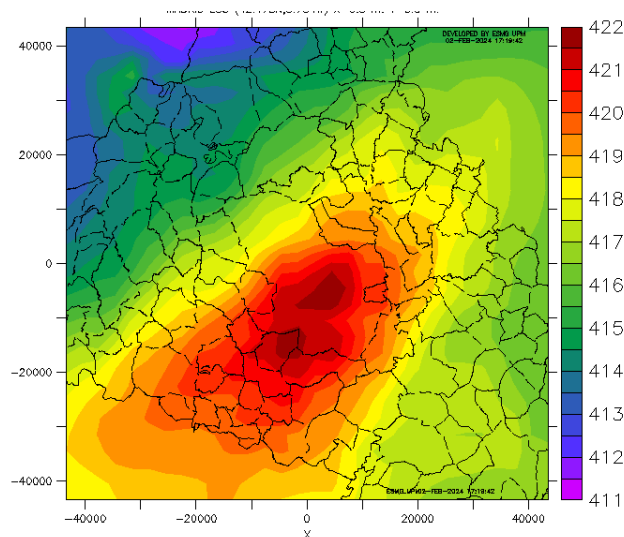
289  
290 Figure 2 shows the differences in expected net biogenic emissions (NEE) between the SSP3-7.0 and SSP5-  
291 8.5 climate scenarios. It shows that in the SSP3-7.0 scenario there would be an increase in biogenic CO<sub>2</sub>

292 emissions in most areas of the Iberian Peninsula. This increase is due to an increase in vegetation  
 293 respiration, which is favoured by the higher temperatures predicted in the SSP3-7.0 scenario compared to  
 294 SSP5-8.5.

295  
 296 Figure 3 provides a visual representation of the spatial distribution of CO<sub>2</sub> concentrations within the 3 km  
 297 spatial resolution domain under the SSP5-8.5 scenario. Concentrations of Figure 3 are the result of a  
 298 combination of anthropogenic and biogenic emissions, boundary conditions, and CO<sub>2</sub> transport and  
 299 diffusion under the simulated meteorological conditions. The figure clearly illustrates the dominance of  
 300 anthropogenic emissions from Madrid within the region. The emitted CO<sub>2</sub> shifts significantly to the  
 301 southwest due to the prevailing winds in the area. In contrast, the mountainous regions have the lowest CO<sub>2</sub>  
 302 concentrations. This phenomenon can be attributed to the presence of sinks in this area, minimal  
 303 anthropogenic emissions and prevailing winds that do not transport CO<sub>2</sub> from the city to these locations.



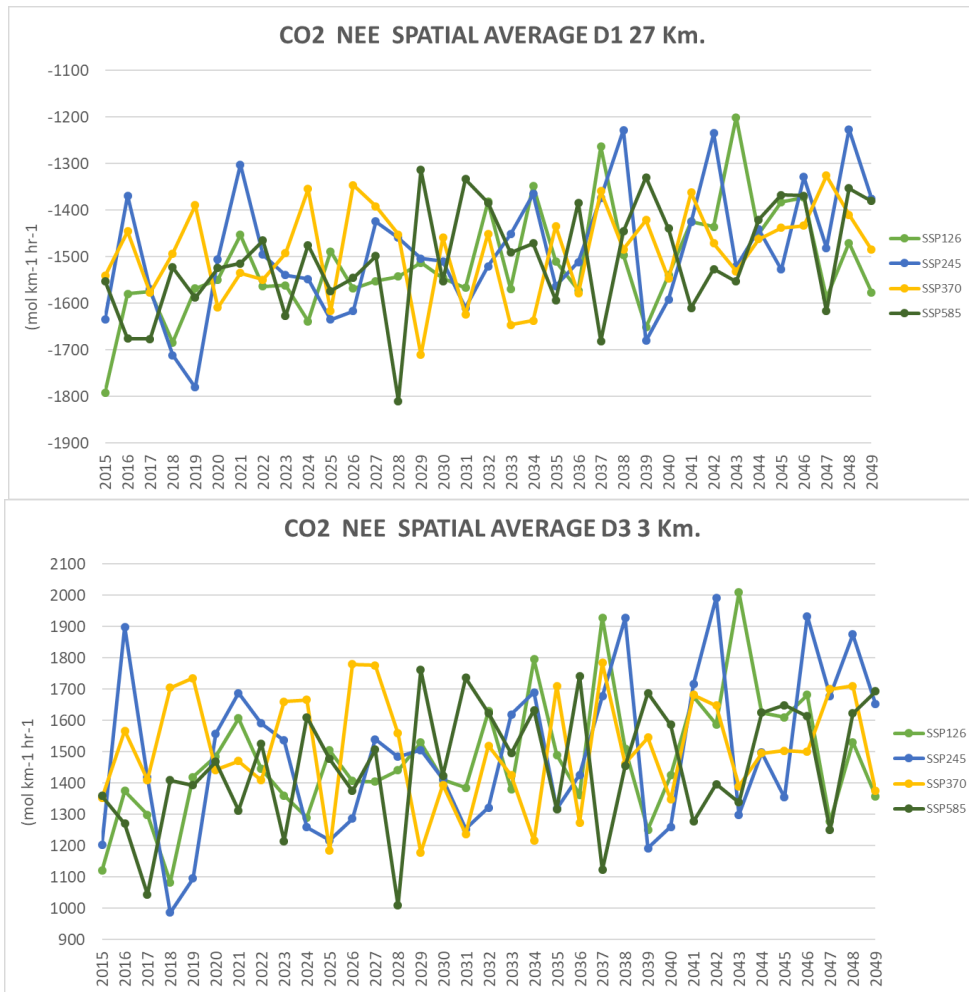
304  
 305 **Figure 2.** Spatial distributions of differences between scenarios SSP370-SSP585 for the NEE (Net  
 306 Ecosystem Exchange) (kmol km<sup>-2</sup> hr<sup>-1</sup>). Mean of 35 years 2015-2050 from WRF/Chem-VPRM  
 307 simulation (27 km).  
 308



309  
 310 **Figure 3.** Spatial distributions of simulated CO<sub>2</sub> concentrations (ppm). Mean of 35 years 2015-2050 from  
 311 WRF/Chem-VPRM simulation (3 km) for SSP585 scenario.  
 312

313 The figure 4 shows the annual evolution of Net Ecosystem Exchange (NEE) of CO<sub>2</sub>, spatially averaged  
 314 over the model domain, for the period 2015–2049 under four CMIP6 scenarios (SSP1-2.6, SSP2-4.5, SSP3-  
 315 7.0, SSP5-8.5). Negative values indicate net CO<sub>2</sub> uptake by vegetation, while positive values represent net  
 316 emissions to the atmosphere  
 317

318



319

320

321

322 **Figure 4.** Temporal evolution of Net Ecosystem Exchange (NEE, CO<sub>2</sub> vegetation flux) [mol km<sup>-1</sup> hr<sup>-1</sup>].  
323 Annual and spatially averaged values over the simulation domain of 27 km domain (D1) and 3km domain  
324 (D3) for each scenario, 2015–2049.

325

326 The figure 4 also reveals spatial differences in the behaviour of biogenic CO<sub>2</sub> fluxes between domains.  
327 Specifically, we observe that in Domain 1 (D1), which covers the entire Iberian Peninsula, net CO<sub>2</sub>  
328 removals dominate over emissions, indicating that, on average, vegetation in this wider region acts as a  
329 carbon sink. In contrast, in Domain 3 (D3) - which covers the Madrid metropolitan area - biogenic CO<sub>2</sub>  
330 emissions outweigh removals. This contrast can be attributed to a combination of vegetation type (in the  
331 north of the Peninsula there is much denser vegetation with higher absorption capacity than in the centre of  
332 the Peninsula), land cover heterogeneity and climatic conditions. In Domain 1, the presence of extensive  
333 forested and semi-natural areas with dense vegetation contributes to higher levels of gross primary  
334 productivity (GPP), thereby increasing net CO<sub>2</sub> uptake. In addition, these areas often experience more  
335 favourable microclimatic conditions (e.g. cooler temperatures and higher soil moisture) for photosynthesis,  
336 especially in elevated or coastal regions. On the other hand, Domain 3 includes urban and peri-urban  
337 landscapes, where vegetation is more fragmented and often consists of grasses, shrubs or ornamental  
338 species with lower photosynthetic efficiency. In addition, urban heat island effects and elevated  
339 temperatures can increase ecosystem respiration (RES), especially during summer months, thereby  
340 counteracting photosynthetic CO<sub>2</sub> uptake and resulting in net positive biogenic emissions.

341

342 When analyzing temporal trends across different climate scenarios, the simulations reveal a progressive  
343 decline in net CO<sub>2</sub> uptakes in Domain 1 (Iberian Peninsula) over time. While this domain maintains its role  
344 as a net carbon sink throughout the simulation period, the magnitude of uptake decreases slightly over time.  
345 This reduction is likely associated with climatic stress on vegetation, including higher temperatures. In  
346 Domain 3 (Community of Madrid), the results show an increase in net biogenic CO<sub>2</sub> emissions over the  
347 same period. Overall, both domains show a deterioration in the biogenic carbon balance. While Domain 1

348 experiences a reduction in its CO<sub>2</sub> sequestration capacity, Domain 3 transforms into a net biogenic CO<sub>2</sub>  
349 source. These trends highlight the sensitivity of biosphere CO<sub>2</sub> fluxes to climate change and suggest that  
350 even natural or semi-natural ecosystems could experience a weakening of their carbon sink function under  
351 future climate conditions.

#### 352 4. CONCLUSIONS

353 In this study, we employed a high-resolution dynamic downscaling system (WRF/Chem-VPRM-EMIMO)  
354 to assess the impact of future climate scenarios on biogenic CO<sub>2</sub> fluxes in the Community of Madrid and  
355 the Iberian Peninsula as a whole. The modeling framework employed a nested approach with three spatial  
356 domains (27 km, 9 km, and 3 km resolution) and focused on the period 2015–2050. Four CMIP6 scenarios  
357 (SSP1-2.6, SSP2-4.5, SSP3-7.0, and SSP5-8.5) were upscaled from the global climate model MPI-ESM1.2-  
358 HR. To isolate climate effects, anthropogenic emissions, land use, and satellite vegetation indices were  
359 fixed at 2018 levels. Model performance was evaluated against hourly temperature observations from three  
360 meteorological stations in the Community of Madrid over the period 2015–2023. The WRF/Chem-VPRM  
361 model showed good agreement with observed data, with R<sup>2</sup> values above 0.80 in all scenarios. Among the  
362 scenarios, SSP5-8.5 presented the best overall fit to observations, although no substantial differences were  
363 found between scenarios in terms of model accuracy. All scenarios projected a gradual increase in average  
364 temperatures over the simulation period, with the largest increase in SSP3-7.0, exceeding SSP5-8.5 by  
365 approximately 0.22%.

366 Spatial analysis of the net CO<sub>2</sub> budget—combining biogenic and anthropogenic fluxes—revealed  
367 contrasting patterns across domains. The northern Iberian Peninsula and parts of the Community of Madrid  
368 function as carbon sinks, primarily due to their extensive vegetation cover. However, in the central  
369 metropolitan area of Madrid, anthropogenic and biogenic emissions predominate, resulting in a net source  
370 of CO<sub>2</sub>. Despite these local emissions, Domain 1 (27 km resolution) still shows an overall net uptake of up  
371 to 1400 mol km<sup>-1</sup> h<sup>-1</sup>, with urban areas such as Madrid contributing positive local fluxes of up to 4500 mol  
372 km<sup>-1</sup> h<sup>-1</sup>. When analyzing the Net Ecosystem Exchange (NEE), we observed that the SSP3-7.0 scenario  
373 presented the largest increase in biogenic CO<sub>2</sub> emissions. This trend is attributed to rising temperatures,  
374 which enhance plant respiration more than photosynthesis, especially in semi-arid urbanized environments.  
375 The other scenarios also showed increases in emissions, albeit to a lesser extent, consistent with their more  
376 moderate temperature trends.

377 In the Community of Madrid, vegetation-related CO<sub>2</sub> emissions exceed absorptions in most areas, with the  
378 exception of mountainous areas, which remain net absorbers due to denser and more productive vegetation.  
379 This spatial heterogeneity underscores the importance of high-resolution modeling in urban and peri-urban  
380 landscapes. Regarding atmospheric CO<sub>2</sub> concentrations, the SSP3-7.0 scenario also showed a slight  
381 increase (~0.11%) relative to SSP5-8.5. On the other hand, the SSP1-2.6 and SSP2-4.5 scenarios projected  
382 small decreases in CO<sub>2</sub> concentrations. However, these variations are moderate, likely due to the strong  
383 influence of boundary conditions on CO<sub>2</sub> concentration fields.

384 As general conclusions of the study, we can highlight that climate change will reduce the capacity of  
385 ecosystems as carbon sinks due to increased respiration and thermal stress, while in urban areas, they will  
386 act as net sources of biogenic CO<sub>2</sub>. High-resolution integrated modeling frameworks, such as WRF/Chem-  
387 VPRM-EMIMO, are essential for capturing spatial heterogeneity and generating valuable information for  
388 the development of mitigation and air quality strategies. Although the study was applied to the Community  
389 of Madrid, this modeling methodology is transferable to other regions and can help decision-makers assess  
390 the impact of local interventions—such as increased vegetation or agroforestry practices—on atmospheric  
391 CO<sub>2</sub> dynamics.

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510

511 **Declarations**

512 **Ethics approval and consent to participate**

513 All authors have followed the ethics rules and have approved to participate.

514 **Consent for publication**

515 All authors have provided the consent to publish this research paper.

516 **Availability of data and material**

517 The datasets generated during and/or analysed during the current study are available in the project website:  
518 <https://distender.eu/>

519 **Competing Interests**

520 The authors have no relevant financial or non-financial interests to disclose

521 **Funding**

522 This work was supported by European Union's Horizon EU research, grant agreement No 101056836

523 **Authors Contributions**

524 All authors have contributed significantly to this work in their respective areas of expertise, ensuring a  
525 comprehensive and multidisciplinary approach to the subject. RSJ prepared the initial draft, outlining the  
526 main concepts, methodologies and conclusions. This draft was revised, expanded and refined by JLP and  
527 MG, who contributed critical ideas, additional data and methodological improvements based on their  
528 specific experience. All authors carefully reviewed the final version of the manuscript, gave their approval,  
529 and agreed to its submission.